

REVISION 1

Structural ordering in calcioferrite-group minerals. Refinement of the crystal structure of fanfaniite from the Hühnerkobel pegmatite, Bavaria.

Ian E. Grey¹, Stephanie Bird², Colin M. MacRae¹, Christian Rewitzer³ and Rupert Hochleitner⁴.

¹CSIRO Mineral Resources, Private Bag 10, Clayton South, Victoria 3169, Australia.

²Australian Synchrotron, 800 Blackburn Road, Clayton, Victoria 3168, Australia.

³Graf von Bogen, Str. 6, 93437 Furth im Wald, Germany.

⁴Mineralogical State Collection (SNSB), Theresienstrasse 41, 80333, München, Germany.

Corresponding author: Ian E. Grey (email: ian.grey@csiro.au)

Abstract

Fanfaniite, $\text{Ca}_4\text{Mn}^{2+}\text{Al}_4(\text{PO}_4)_6(\text{OH})_4 \cdot 12\text{H}_2\text{O}$, from the Hühnerkobel pegmatite mine, Bavaria, has been characterized by chemical analyses and synchrotron single-crystal diffraction. The average crystal structure was refined in space group $C2/c$ (cell parameters $a = 10.055(2)$, $b = 24.132(5)$, $c = 6.2590(10)$ Å, $\beta = 91.35(3)^\circ$) to compare with reported monoclinic structures of other calcioferrite-group minerals with general formula $\text{Ca}_4\text{AB}_4(\text{PO}_4)_6(\text{OH})_4 \cdot 12\text{H}_2\text{O}$, $A = \text{Mn}^{2+}$, Fe^{2+} , Mg , $B = \text{Al}$, Fe^{3+} . The average structure contains disordered half-occupied A sites and associated coordinated water molecules. The diffraction data for fanfaniite contains weak reflections that violate the c -glide condition, as also reported for montgomeryite, and in addition contains extremely weak, diffuse reflections requiring a doubling of a , as reported for kingsmountite. Structure refinements were conducted for the noncentrosymmetric $C2$ model used for montgomeryite and for the $P-1$ model used for kingsmountite. The fanfaniite diffraction data is better explained by the triclinic model with doubled a cell parameter, although the extent of ordering of the A-site cations is considerably lower (56%) than reported for kingsmountite (85%). If the $C2$ model contributes, it can only be at the scale of the unit cell.

1 Introduction

The calcioferrite group of minerals has the general formula $\text{Ca}_4\text{A}^{2+}\text{B}^{3+}_4(\text{PO}_4)_6(\text{OH})_4 \cdot 12\text{H}_2\text{O}$, where $\text{A} = \text{Mg}, \text{Mn}^{2+}$ and Fe^{2+} and $\text{B} = \text{Al}$ and Fe^{3+} (Grey et al., 2019a). Although calcioferrite, with $\text{A} = \text{Mg}, \text{B} = \text{Fe}^{3+}$, was the first member to be characterised (Blum, 1858), montgomeryite, with $\text{A} = \text{Mg}, \text{B} = \text{Al}$, was the first member to have the crystal structure determined, in space group $C2/c$ with cell parameters $a = 10.023(1)$, $b = 24.121(3)$, $c = 6.243(1)$ Å, $\beta = 91.55(1)^\circ$ (Moore and Araki, 1974). The structure is built from [100] zig-zag chains of alternating *cis*- and *trans*- corner-shared octahedra that share corners with PO_4 tetrahedra along [001] to form heteropolyhedral slabs parallel to (010). The larger Ca and A-site atoms occupy sites at the periphery of, and between, the slabs. Studies on the other members, kingsmountite, with $\text{A} = \text{Fe}^{2+}, \text{B} = \text{Al}$ (Dunn et al. 1979), zodacite, $\text{A} = \text{Mn}^{2+}, \text{B} = \text{Fe}^{3+}$ (Dunn et al., 1988; Lafuente et al., 2014) and fanfaniite, with $\text{A} = \text{Mn}^{2+}, \text{B} = \text{Al}$ (Grey et al., 2019b) all reported the *C*-centred monoclinic cell with similar unit-cell parameters as for montgomeryite.

A feature of the monoclinic $C2/c$ structure is disordering of the A site atoms, with pairs of edge-shared A-centred octahedra that are only half-occupied. This aspect was addressed by Fanfani et al. (1976). In their single-crystal diffraction study on montgomeryite, they observed weak reflections that violated the *c*-glide extinction condition. They noted that the reflection conditions were consistent with space group $C2$, and that the A site split into two sites in $C2$, allowing ordering at one of these sites. They were unable to refine the whole structure in $C2$ because of high correlations between the positional and thermal parameters of related atoms, but a refinement restricted to the A site cation and coordinated H_2O resulted in a lowering of the *R* factor.

A further modification to the structural ordering in calcioferrite minerals was found in a study of a kingsmountite-related member from the Foote mine, North Carolina (Grey et al., 2019a). A structure analysis showed it to have a triclinic superstructure of the monoclinic structure, with a doubled a_m cell parameter. Refinement of the crystal structure showed that the superstructure was due to ordering of the A site cations, as well as ordering of octahedrally-coordinated Mn at one of the Ca sites, giving an ideal formula

$\text{Ca}_3\text{Mn}^{2+}\text{Fe}^{2+}\text{Al}_4(\text{PO}_4)_6(\text{OH})_4 \cdot 12\text{H}_2\text{O}$. A study was then conducted on the type specimen of kingsmountite, which showed that it also had a doubling of the a_m cell parameter, resulting in a redefinition of the ideal formula and space group for kingsmountite (Grey et al., 2019a).

During the course of a study of secondary phosphate minerals from the Hühnerkobel pegmatite mine, Bavaria, we found well developed crystals of a calcioferrite-group mineral

that was subsequently analysed as being fanfaniite. A synchrotron single-crystal diffraction data collection on one of the crystals was indexed by the same monoclinic cell as reported for type fanfannite and a good refinement ($R_{\text{obs}} = 0.028$ for 1989 observed reflections) was obtained in space group C2/c. The data analysis, however, showed the presence of 38 reflections that violated the c -glide condition. Although weak, the reflections had intensities, I , up to $10\sigma(I)$, and could not be ignored. In addition, during the data collection, screen-shots of the diffraction frames showed very weak diffuse reflections corresponding to a doubling of the a_m cell parameter. These reflections, which were streaked parallel to a^* , were not picked up by the synchrotron automated indexing routine, but we were able to have them incorporated by reprocessing the frame data with forcing of the superstructure unit cell found for kingsmountite. The data analysis showed that there were more than 600 such reflections with $I > 3\sigma(I)$ in the reprocessed dataset.

The different datasets collected on the same crystal provided an opportunity to compare the analysis of structural order based on different models, the results of which are reported here.

2 Specimen description and chemical analysis

The fanfaniite-containing specimens were collected from the dumps of the Hühnerkobel pegmatite mine, near the town of Zwiesel, Bavaria. The mineral is found in cavities of K-feldspar with traces of apatite, rockbridgeite, and mitridatite. Associated minerals are jahnsite-CaMnFe, kayrobertsonite, and the paulkerite-group minerals rewtizerite and macraeite. Fanfaniite forms sprays of transparent, light yellow blades, shown in Fig. 1. The blades are up to 120 μm in length and 30 μm wide.

Crystals of fanfaniite, mounted in epoxy blocks and polished, were analysed using wavelength-dispersive spectrometry on a JEOL JXA 8530F Hyperprobe operated at an accelerating voltage of 15 kV and a beam current of 2 nA. The beam was defocused to 2 to 10 μm , depending on the size of the crystal. To minimise dehydration during analysis a cold stage cooled to liquid nitrogen temperature was employed in the microprobe, and the specimen was precooled under dry nitrogen prior to introduction into the microprobe. Analytical results (average of 10 analyses), are given in Table 1. There was insufficient material for direct determination of H_2O , so it was based upon the crystal structure, with 14 H_2O per 6 P. The calculated H_2O content was included in the in-house matrix correction procedure. Bond valence sum (BVS) values calculated from the refined M-O distances were consistent with all iron as Fe^{3+} .

The empirical formula, in structural form, scaled to 15 (M+P) is:

$(\text{Ca}_{3.84}\text{Mn}^{2+}_{0.18})(\text{Mn}^{2+}_{0.48}\text{Mg}_{0.21}\text{Fe}^{3+}_{0.31})(\text{Al}_{3.36}\text{Fe}^{3+}_{0.64})(\text{PO}_4)_{5.97}[(\text{OH})_{3.56}\text{O}_{0.44}]\cdot 12.15\text{H}_2\text{O}$. The ideal formula with dominant cations at each site is $\text{Ca}_4\text{Mn}^{2+}\text{Al}_4(\text{PO}_4)_6(\text{OH})_4\cdot 12\text{H}_2\text{O}$, which corresponds to fanfaniite.

3 Crystallography

A fanfaniite blade, measuring 0.090 x 0.030 x 0.010 mm, was used for a data collection at the Australian Synchrotron microfocus beamline MX2 (Aragao *et al.*, 2018). Intensity data were collected using a Dectris Eiger 16M detector and monochromatic radiation with a wavelength of 0.7107 Å. The crystal was maintained at 100 K in an open-flow nitrogen cryostream during data collection. The diffraction data were collected using a single 36 second sweep of 360° rotation around phi. The resulting dataset consists of 3600 individual images with an approximate phi angle of each image being 0.1 degrees. The raw intensity dataset was processed using CrysAlisPro software (Agilent, 2014) to produce data files that were analysed using SHELXT (Sheldrick, 2015) and JANA2006 (Petříček *et al.*, 2014).

3.1 Average C2/c structure

Automatic indexing of the frame diffraction data gave a C-centred monoclinic unit cell with parameters given in Table 2. A structure solution in space group C2/c was obtained using SHELXT. It was found to be the same that reported for fanfaniite from the Foote mine (Grey *et al.*, 2019). For the refinement, the scattering curve for Mn was used for the A site and the site occupancy was refined, while Fe and Al were assigned to the Al sites and their relative amounts were refined. Refinement in JANA2006 with anisotropic displacement parameters showed that one of the water molecules coordinated to Mn, Ow3, had very large U_{22} and U_{33} parameters (0.07 Å^2), indicating a likely split Ow3, as reported also for montgomeryite (Fanfani *et al.*, 1976) and for calcioferrite (Lafuente *et al.*, 2014). Splitting of Ow3 into Ow3a and Ow3b and independent refinement of the atoms at the two sites resulted in sensible anisotropic ADP values of $\sim 0.02 \text{ Å}^2$. Difference Fourier maps showed unambiguous locations for H atoms associated with Ow1, Ow2 and OH. These were included, without refinement of the coordinates and with refinement of a common isotropic ADP. The final refinement, with anisotropic ADPs for all non-H atoms, converged to $R_{\text{obs}} = 0.028$ for 1989 observed reflections with $I > 3\sigma(I)$. Other details of the refinement are given in Table 2. The refined coordinates, equivalent ADPs and site occupancies are reported in Table 3, together with BVS values obtained using the parameters of Gagné and Hawthorne (2015). The BVS values for the Mn sites were calculated using the site occupancies in the empirical formula, $\text{Mn}_{0.48}\text{Al}_{0.21}\text{Fe}_{0.31}$. The resulting value, 2.35, is consistent with all Fe as Fe^{3+} , which gives a formal BVS of 2.31, and suggests that a small amount of Mn may be present as Mn^{3+} .

The polyhedral bond distances are reported in Table 4. An interesting result from the refinement is that Fe is preferentially ordered in the Al2 site, which contains 36% Fe compared with only 12% Fe at the Al1 site. As shown in Fig. 2(a), the Al1-centred octahedron shares two edges with Ca2O₈ polyhedra whereas the Al2 octahedron is involved in only corner-sharing with other polyhedra, and so has more flexibility to accommodate a larger cation.

3.2 C2 refinement

The processing of the fanfaniite dataset using space group C2/c gave 38 merged reflections that violated the *c*-glide extinction condition ($h0l, l=2n+1$) of which 15 had $I > 3\sigma(I)$. Although very weak, the presence of these reflections indicate that fanfaniite from Hühnerkobel has a symmetry lower than C2/c. There are two maximal non-isomorphic subgroups of C2/c. These are C2 and P-1. Only the first of these allows splitting of the A site into two sites with the possibility of ordering of the A site cations in one of the sites. A similar situation of the presence of weak reflections violating the *c*-glide condition was reported by Fanfani et al. (1976) for montgomeryite. These authors proposed that Mg was ordered into one of the split sites in a C2 modelling of the structure, although they were not able to refine the complete C2 model because of strong correlations between the parameters of the split sites.

The C2/c coordinate and reflection files were transformed into the C2 subgroup using JANA2006. The site occupancies at the split Mn1 and Mn2 sites were first refined, and then the coordinates and anisotropic ADPs were refined with restrictions to keep equal the ADPs of the split Mn and Ow3 site atoms. H atoms coordinated to oxygen at the OH, Ow1 and Ow2 pairs of sites were located in difference Fourier maps and included in the refinement with the coordinates fixed and with refinement of a common isotropic ADP. The refinement converged to $R_{obs} = 0.028$ for 3890 observed reflections with $I > 3\sigma(I)$. The partial R_{obs} for the 15 observed $h0l, l = 2n+1$ reflections was 0.089. Other refinement results are given in Table 2. The refined parameters, including calculated BVS values are reported in Table 5 and the polyhedral bond distances are given in Table 6. Although there were numerous high correlations between the parameters of split pairs of atoms in the C2 refinement, the crystal chemistry of the refined structure remained sound, as indicated by the normal ranges of bond distances in Table 6 and the BVS values very close to those obtained in the C2/c refinement. The refined site occupancies at the Mn sites were 0.510(3) for Mn1 and 0.378(3) for Mn2, with the mean value of 0.444 matching the refined Mn site occupancy of 0.445(2) from the C2/c refinement. The C2 model has a small but significant preference of Mn to occupy the

Mn1 site. Atoms at the Mn1 site coordinate to Ow3a2 and Ow3b2 (Table 6), and it would be expected that the refined site occupancies at these sites should be higher than for Ow3a1 and Ow3b1 that coordinate to Mn2. The refined occupancies at the Ow3 sites in Table 5 do not support this, although the deviations of the site occupancies from 0.5 for all four Ow3 sites are $< 3\text{SDs}$.

As a check on the contribution of the Mn ordering to the $h0l$, $l = 2n+1$ reflections, the site occupancies for the Mn1 and Mn2 sites from the above refinement were both set to the mean value of 0.444 and fixed. Refinement of the rest of the parameters converged at $R_{\text{obs}} = 0.029$, and the partial R_{obs} factor for the observed $h0l$, $l = 2n+1$ reflections was 0.14. On release of the site occupancies at Mn1 and Mn2, subsequent refinement did not significantly change the R factors or site occupancies. Thus the fitting of the “forbidden” c -glide reflections using the $C2$ model can be largely achieved without invoking any order at the A sites.

3.3 $P-1$ refinement with doubled a cell parameter – kinsmountite model

As mentioned in the Introduction, screen images of the diffraction data obtained during the data collection showed rows of strong sharp spots with very weak, streaked spots halfway between. A screenshot showing the diffuse reflections halfway between sharp spots oriented along a^* is shown in Fig. 3. These weak diffuse reflections were not picked up by the synchrotron software in the data indexing routine. By analogy with kinsmountite, the diffuse reflections correspond to a doubling of the a axis. In a transmission electron microscopy study of the holotype kinsmountite mineral (Grey et al., 2019a), we observed weak diffuse reflections corresponding to a doubling of the a axis and we obtained a good refinement of the structure for a kinsmountite crystal from the Foote mine using a triclinic unit cell, related to the monoclinic cell by the matrix $\begin{pmatrix} 2 & 0 & 0 \\ 0 & -0.5 & 0.5 \\ 0 & 0 & 1 \end{pmatrix}$, with a doubling of a . The refinement confirmed 85% ordering of the A-site cation (Fe^{2+}) in kinsmountite.

The synchrotron frame data for fanfaniite was reprocessed by forcing a kinsmountite-type triclinic cell with doubled a parameter. The resulting reflection file contained an additional 3822 reflections with h odd, of which 17% (656 reflections) had $I > 3\sigma(I)$. The refinement input file for triclinic kinsmountite was used to initiate the refinement in $P-1$, with Mn replacing Fe at the two A sites and with Ca replacing Mn at the Mn3 site. Mixed Al/Fe occupancy was used for the 5 independent Al sites. The Ow1 to Ow4 sites, involved in coordination to the A site cations, were each split into pairs, as was done for Ow3 in the $C2/c$ refinement. After cycles of anisotropic ADP refinements, a difference Fourier map was used to locate 25 of the possible 28 H atoms. These were included and a H group isotropic ADP was refined. The final refinement converged at $R_{\text{obs}} = 0.025$ for 4266 observed reflections

with $I > 3\sigma(I)$. An R_{obs} value of 0.19 was obtained for 656 observed reflections with $h = 2n+1$. Although this is elevated, it is not unreasonable considering that most of the observed reflections giving the doubled a cell parameter are barely above $3\sigma(I)$. The other refinement details are given in Table 2. The refined coordinates, equivalent isotropic ADPs, site occupancies and BVS values are reported in Table 7 and polyhedral bond distances in Table 8.

The site occupancies for the split A sites, based on using the Mn scattering curve for the mixed $\text{Mn}_{0.48}\text{Al}_{0.21}\text{Fe}_{0.31}$ occupancy, are 0.497(1) for Mn1 and 0.397(1) for Mn2. Scaling the summed 0.894 occupancy to 1 gives values of 0.556 and 0.444 for the corrected occupancies. It is interesting to note that the refined site occupancies for Ow1a to Ow4a, coordinated to Mn1 are in the ranges 0.56 to 0.61 and those for Ow1b to Ow4b, coordinated to Mn2 are in the range 0.39 to 0.44, which are reasonably consistent with the occupancies of the metal sites. This is in contrast to the $C2$ refinement where there was no relationship between the Mn site occupancies and the occupancies of the coordinating water molecules.

The H-bonding scheme for the kingsmountite P -1 model of fanfaniite, is given in Table 9. The oxygen atoms that have low BVS values in Table 7 (calculated for metal atom contributions only) all feature as acceptor atoms for H bonds in Table 9. Noteworthy are the strong H-bonds, with $D-A \sim 2.60$ Å, associated with Ow5, Ow6, Ow11 and Ow12 as donors to oxygens shared between Al and P. The H bond between Ow3a and O22 is associated with a particularly short $D-A$ distance of 1.53 Å. A similar short $D-A$ distance of 1.54 Å involving Ow3 has also been reported for montgomeryite (Fanfaniite et al., 1976).

4. Discussion

The $C2/c$ structural model that has been widely reported for calcioferrite group minerals is an average structure, with disordered, half-occupied A sites and associated coordinated water molecules. Two alternative models have been reported to describe ordering of the A site cations: the $C2$ model proposed for montgomeryite by Fanfani et al. (1976) and the P -1 model with doubled a cell parameter reported for kingsmountite (Grey et al., 2019a).

Although the non-centrosymmetric $C2$ model gave a satisfactory refinement for fanfaniite, as indicated by the results in Tables 2, 5 and 6, there are some concerns as to its validity. An analysis of the normalised structure factors (E values) for the fanfaniite diffraction data gave a very good fit to the distribution curve for a centrosymmetric model. The mean value of $|E^2 - 1|$ of 0.939 compares with values of 0.968 for centrosymmetry and 0.736 for non-centrosymmetry, giving a probability of the structure being centrosymmetric of 90.2%. If the $C2$ model is valid, it can only be so for extremely small regions, of a size comparable to the

unit cell (nanoscale domains). A refinement of the *C2* model as a twinned structure with the inversion centre as the twin operator gave exactly 50% for the volumes of the two twin domains. This is another indication that the twin domains must be on a very fine scale. A projection of the *C2* model along [001] is shown in Fig. 2(b). The octahedron is plotted only for the more-highly occupied *A* site, Mn1. A comparison with the average structure in Fig. 2(a) shows that the ordering at the *A* sites in *C2* gives occupied *A*-centred octahedra on one side only of the heteropolyhedral layers.

The [001] projection of the kingsmountite-type triclinic model with doubled *a* cell parameter is shown in Fig. 4. In contrast to the *C2* model, the *P*-1 doubled-*a* model has ordering of the *A*-site cations on both sides of the heteropolyhedral layers, and with an alternation of filled and vacant sites (for complete order) along [100]. Whereas for kingsmountite the *A*-site order was 85%, for fanfaniite from Hühnerkobel the ordering is only 56%. Such a low degree of ordering was manifested in extremely weak and diffuse reflections with $h=2n+1$, that were streaked along a^* .

In kingsmountite, in addition to the *A*-site ordering, there is also partial substitution of octahedrally coordinated Mn for 8-coordinated Ca at the Ca3 site (9% Mn at Mn3) and the Ca4 site (82% Mn at Mn4). The Ca site composition for kingsmountite is $\text{Ca}_{3.15}\text{Mn}_{0.85}$, corresponding to 21% substitution of Mn. The fanfaniite empirical composition contains only a very minor 4 at% substitution of Mn at the Ca sites. Difference Fourier maps from the fanfaniite refinement showed very weak peaks ($<0.8 \text{ e.Å}^3$) displaced by $\sim 0.8 \text{ Å}$ from Ca3 and Ca4 that corresponded to the Mn3 and Mn4 positions in kingsmountite. Inclusion of these atoms in the refinement gave site occupancies of only 0.01-0.02, so they were not incorporated into the final refinement. Nevertheless, their presence further confirms the validity of the kingsmountite triclinic structure for fanfaniite.

In conclusion, the triclinic structure has been validated for fanfaniite from Hühnerkobel, while there are doubts about the *C2* model; if it contributes, then it does so in very small, unit-cell-scale domains. Future studies on new and known calcioferrite-group minerals should examine very closely the diffraction patterns for possible doubling of the *a* axis relative to the commonly determined *C2/c* average structure unit cell, to check if the triclinic model ordering is applicable.

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Table 1. Chemical data (wt%) for fanfaniite from Hühnerkobel.

Constituent	Mean	Range	SD	Standard
CaO	17.79	16.14-18.79	0.86	Wollastonite
MnO	3.84	2.99-4.76	0.65	MnSiO ₃
MgO	0.70	0.55-0.96	0.12	Spinel
Fe ₂ O ₃	6.30	4.31-7.97	1.26	Hematite
Al ₂ O ₃	14.14	12.82-15.13	0.74	Berlinite
P ₂ O ₅	34.99	32.68-37.19	1.26	Berlinite
H ₂ O _{calc}	20.71			
Total	98.47			

Table 2. Crystal data and structure refinement results.

Space group	<i>C2/c</i>	<i>C2</i>	<i>P</i> -1
Unit cell dimensions, Å, °.	<i>a</i> = 10.055(2) <i>b</i> = 24.132(5) <i>c</i> = 6.2590(10) β = 91.35(3)	<i>a</i> = 10.055(2) <i>b</i> = 24.132(5) <i>c</i> = 6.2590(10) β = 91.35(3)	<i>a</i> = 20.0730(1) <i>b</i> = 13.0393(7) <i>c</i> = 6.2483(4) α = 89.462(5) β = 91.3919(5) γ = 112.629(5)
Volume, Å ³ , <i>Z</i>	1518.3(5), 2	1518.3(5), 2	1508.84(16), 2
Temperature (K)	100	100	100
Density (calc.), g/cm ³	2.61	2.61	2.62
Crystal size, mm	0.090 x 0.030 x 0.010		
Absorption coefficient, mm ⁻¹	2.07	2.07	2.06
SADABS, <i>T</i> _{min} / <i>T</i> _{max}	0.65, 0.75		
Reflections collected, <i>R</i> _{int}	13,207, 0.049	13,338, 0.047	25,088, 0.022
Reflection range	14 < <i>h</i> < 14 -33 < <i>k</i> < 33 -9 < <i>l</i> < 9	-14 < <i>h</i> < 14 -33 < <i>k</i> < 33 -9 < <i>l</i> < 9	-29 < <i>h</i> < 29 -17 < <i>k</i> < 17 -9 < <i>l</i> < 9
Unique refl. - all, <i>I</i> > 3σ(<i>I</i>)	2099, 1989	4199, 3890	7556, 4266
Number refined parameters	144	264	510

Final R indices [$I > 3\sigma(I)$]	$R_{\text{obs}} = 0.028$ $wR_{\text{obs}} = 0.038$	$R_{\text{obs}} = 0.028$ $wR_{\text{obs}} = 0.038$	$R_{\text{obs}} = 0.025$ $wR_{\text{obs}} = 0.041$
R indices (all data)	$*R_{\text{obs}} = 0.030$ $wR_{\text{obs}} = 0.039$	$R_{\text{obs}} = 0.031$ $wR_{\text{obs}} = 0.038$	$R_{\text{obs}} = 0.041$ $wR_{\text{obs}} = 0.045$
GofF	2.28	1.81	2.20
Largest diff. peak, hole ($\text{e}\text{\AA}^{-3}$)	0.91, -0.53	0.86, -0.58	0.86, -0.54

$$*w = [\sigma^2(F_o) + (0.01F_o)^2]^{-1}$$

Table 3. Coordinates, site occupancies, equivalent isotropic displacement parameters and bond valence sums (valence units) for C2/c model for fanfaniite

	occupancy	x	y	z	$U_{\text{eq}} (\text{\AA}^2)$	BVS
Ca1	1	0	0.061296(16)	0.25	0.01013(11)	1.80
Ca2	1	0	0.330446(16)	0.25	0.00842(11)	2.10
Al1	0.876(4)	0.25	0.25	0	0.00624(15)	3.17
Fe1	0.124(4)	0.25	0.25	0	0.00624(15)	
Al2	0.638(4)	0	0.170374(16)	0.75	0.00646(13)	3.04
Fe2	0.362(4)	0	0.170374(16)	0.75	0.00646(13)	
Mn	0.445(2)	0	0.47269(3)	0.25	0.0089(2)	2.35
P1	1	0.5	0.30049(2)	0.75	0.00699(13)	5.00
P2	1	0.26051(3)	0.115500(14)	0.96241(6)	0.00918(10)	4.94
O1	1	0.61661(10)	0.26145(4)	0.70566(17)	0.0104(2)	2.00
O2	1	0.46901(10)	0.33818(4)	0.55756(18)	0.0114(3)	1.86
O3	1	0.31089(10)	0.17509(4)	0.00908(17)	0.0107(3)	1.94
O4	1	0.37755(11)	0.08723(5)	0.85518(19)	0.0134(3)	1.77
O5	1	0.13937(10)	0.11642(4)	0.80334(19)	0.0136(3)	1.73
O6	1	0.22078(13)	0.08710(5)	0.1691(2)	0.0212(3)	1.60
Oh1	1	0.36898(10)	0.27120(4)	0.22367(16)	0.0091(2)	1.06
Ow1	1	0.16539(11)	0.32776(5)	0.5213(2)	0.0184(3)	0.34
Ow2	1	0.11229(11)	0.02551(4)	0.58020(17)	0.0134(3)	0.42
Ow3a	0.5	0.1186(5)	0.46599(16)	0.5971(6)	0.0198(10)	0.25
Ow3b	0.5	0.1249(4)	0.48041(15)	0.5392(6)	0.0148(9)	0.31
Hoh	1	0.3638	0.2443	0.3317	0.063(5)	
H1a	1	0.1708	0.349	0.6014	0.063(5)	
H1b	1	0.2393	0.3093	0.4101	0.063(5)	
H2a	1	0.1061	0.054	0.6494	0.063(5)	
H2b	1	0.200	0.016	0.5567	0.063(5)	

Table 4. Polyhedral distances for C2/c model for fanfaniite

Ca1-O2	x2	2.7230(11)	Ca2-O1	x2	2.5270(11)
Ca1-O6	x2	2.3718(13)	Ca2-O3	x2	2.5315(11)
Ca1-Ow2	x2	2.4863(11)	Ca2-O4	x2	2.4360(11)
Ca1-Ow2	x2	2.6171(11)	Ca2-Ow1	x2	2.3493(12)
Av.		2.549	Av.		2.461
Al1-O1	x2	1.9009(10)	Al2-O2	x2	1.9682(11)
Al1-O3	x2	1.9091(10)	Al2-O5	x2	1.9361(11)
Al1-Oh1	x2	1.8904(10)	Al2-Oh1	x2	1.9341(10)
Av.		1.900	Av.		1.946
Mn-O4	x2	2.0196(12)	P2-O3		1.5501(10)
Mn-Ow3a	x2	2.140(4)	P2-O4		1.5290(12)
Mn-Ow3b	x2	2.163(4)	P2-O5		1.5551(11)
Av.		2.108	P2-O6		1.5254(13)
			Av.		1.540
P1-O1	x2	1.5345(11)			
P1-O2	x2	1.5353(11)			
Av.		1.535			

Table 5. Coordinates, site occupancies, equivalent isotropic displacement parameters and bond valence sums (valence units) for C2 model for fanfaniite.

	occupancy	x	y	z	Ueq (Å ²)	BVS
Ca1_1	1	0	0.06131(6)	0.5	0.0102(3)	1.77
Ca1_2	1	0	-0.06129(6)	0	0.0100(3)	1.83
Ca2_1	1	0	0.33031(6)	0.5	0.0080(3)	2.11
Ca2_2	1	0	-0.33058(6)	0	0.0086(3)	2.08
Al1	0.883(3)	0.24971(14)	0.24964(7)	0.2486(3)	0.00582(12)	3.15
Fe1	0.117(3)	0.24971(14)	0.24964(7)	0.2486(3)	0.00582(12)	
Al2_1	0.671(7)	0	0.17045(5)	1	0.0071(4)	3.07
Fe2_1	0.329(7)	0	0.17045(5)	1	0.0071(4)	
Al2_2	0.579(7)	0	-0.17031(5)	-0.5	0.0063(3)	3.02
Fe2_2	0.421(7)	0	-0.17031(5)	-0.5	0.0063(3)	
Mn_1	0.510(3)	0	0.47263(8)	0.5	0.00870(15)	2.31
Mn_2	0.378(3)	0	-0.47271(11)	0	0.00870(15)	2.41
P1_1	1	0.5	0.30044(8)	1	0.0069(4)	5.00
P1_2	1	-0.5	-0.30054(8)	-0.5	0.0068(4)	5.01
P2_1	1	0.25940(10)	0.11548(5)	1.21147(19)	0.0089(3)	4.97
P2_2	1	-0.26162(11)	-0.11548(5)	-0.71348(20)	0.0093(3)	4.95
O1_1	1	0.6171(3)	0.26083(15)	0.9556(5)	0.0104(8)	1.98
O1_2	1	-0.6157(3)	-0.26214(15)	-0.4561(5)	0.0097(8)	2.01
O2_1	1	0.4676(2)	0.33852(14)	0.8091(4)	0.0101(8)	1.87
O2_2	1	-0.4707(3)	-0.33796(15)	-0.3061(4)	0.0122(8)	1.86

O3_1	1	0.3097(3)	0.17461(14)	0.2603(4)	0.0122(8)	1.94
O3_2	1	-0.3117(3)	-0.17550(13)	0.2414(4)	0.0093(8)	1.93
O4_1	1	0.3779(3)	0.08702(16)	1.1035(5)	0.0158(9)	1.72
O4_2	1	-0.3775(3)	-0.08737(15)	-0.6063(5)	0.0111(9)	1.82
O5_1	1	0.1381(3)	0.11678(14)	1.0488(5)	0.0133(8)	1.73
O5_2	1	-0.1407(3)	-0.11585(15)	-0.5583(5)	0.0135(8)	1.75
O6_1	1	0.2225(3)	0.08691(17)	0.4157(5)	0.0207(10)	1.63
O6_2	1	-0.2192(3)	-0.08741(17)	0.0783(5)	0.0203(10)	1.59
Oh1_1	1	0.3697(3)	0.27152(14)	0.4737(5)	0.0090(8)	1.07
Oh1_2	1	-0.3681(2)	-0.27089(15)	0.0262(5)	0.0091(8)	1.06
Ow1_1	1	0.1659(3)	0.32769(15)	0.7694(5)	0.0175(9)	0.34
Ow1_2	1	-0.1646(3)	-0.32817(18)	-0.2736(5)	0.0195(10)	0.34
Ow2_1	1	0.1124(3)	0.02552(16)	0.8317(5)	0.0139(8)	0.42
Ow2_2	1	-0.1123(3)	-0.02570(15)	-0.3284(5)	0.0133(8)	0.42
Ow3a_1	0.505(18)	0.1175(7)	0.4685(4)	0.8464(19)	0.0190(7)	0.26
Ow3a_2	0.507(19)	-0.1201(7)	-0.4635(4)	-0.3491(18)	0.0190(7)	0.24
Ow3b_1	0.555(19)	0.1249(7)	0.4815(5)	0.7832(14)	0.0190(7)	0.31
Ow3b_2	0.521(18)	-0.1257(7)	-0.4793(5)	-0.2953(14)	0.0190(7)	0.32
Hoh1	1	0.3569	0.2443	0.5891	0.054(3)	
Hoh2	1	-0.3622	-0.244	-0.0931	0.054(3)	
H1a1	1	0.178	0.3481	0.868	0.054(3)	
H1b1	1	-0.166	-0.3494	-0.3559	0.054(3)	
H1b2	1	-0.2425	-0.3113	-0.1755	0.054(3)	
H2a1	1	0.1953	0.0156	0.7959	0.054(3)	
H2a2	1	0.1016	0.056	0.9135	0.054(3)	
H2b1	1	-0.1014	-0.0545	-0.4134	0.054(3)	
H2b2	1	-0.1983	-0.0142	-0.3126	0.054(3)	

Table 6. Polyhedral distances for C2 model for fanfaniite

Mn_1-O4_2	x2	2.023(3)	Mn_2-O4_1	x2	2.011(4)
Mn_1-Ow3a_2	x2	2.186(10)	Mn_2-Ow3a_1	x2	2.094(10)
Mn_1-Ow3b_2	x2	2.158(9)	Mn_2-Ow3b_1	x2	2.173(9)
Av.		2.122			2.093
Ca1_1-O2_2	x2	2.730(4)	Ca1_2-O2_1	x2	2.713(4)
Ca1_1-O6_1	x2	2.392(3)	Ca1_2-O6_2	x2	2.355(3)
Ca1_1-Ow2_1	x2	2.495(3)	Ca1_2-Ow2_1	x2	2.614(4)
Ca1_1-Ow2_2	x2	2.626(4)	Ca1_2-Ow2_2	x2	2.476(3)
Av.		2.561	Av.		2.540
Ca2_1-O1_2	x2	2.534(4)	Ca2_2-O1_1	x2	2.519(4)
Ca2_1-O3_2	x2	2.523(3)	Ca2_2-O3_1	x2	2.545(3)
Ca2_1-O4_2	x2	2.438(4)	Ca2_2-O4_1	x2	2.433(4)
Ca2_1-Ow1_1	x2	2.345(3)	Ca2_2-Ow1_2	x2	2.354(3)
Av.		2.460	Av.		2.463
Al1-O1_1		1.892(3)	Al2_1-O2_2	x2	1.956(3)
Al1-O1_2		1.915(3)	Al2_1-O5_1	x2	1.918(3)
Al1-O3_1		1.910(4)	Al2_1-Oh1_2	x2	1.944(3)

Al1-O3_2	1.909(4)	Av.	1.939
Al1-Oh1_1	1.908(3)		
Al1-Oh1_2	1.874(3)	Al2_2-O2_1 x2	1.981(3)
Av.	1.901	Al2_2-O5_2 x2	1.959(3)
		Al2_2-Oh1_1 x2	1.925(3)
		Av.	1.955
P1_1-O1_1 x2	1.547(3)	P1_2-O1_2 x2	1.517(3)
P1_1-O2_1 x2	1.536(3)	P1_2-O2_2 x2	1.536(3)
Av.	1.541		1.526
P2_1-O3_1	1.542(4)	P2_2-O3_2	1.557(3)
P2_1-O4_1	1.544(4)	P2_2-O4_2	1.518(3)
P2_1-O5_1	1.570(3)	P2_2-O5_2	1.539(3)
P2_1-O6_1	1.506(4)	P2_2-O6_2	1.538(4)
Av.	1.540	Av.	1.538

Table 7. Coordinates, site occupancies, equivalent isotropic displacement parameters and bond valence sums (valence units) for kingmountite triclinic model for fanfaniite.

	occupancy	x	y	z	Ueq (Å ²)	BVS
Mn1	0.497(1)	0.23644(3)	-0.05458(5)	0.75026(11)	0.00797(12)	2.34
Mn2	0.397(1)	0.73613(4)	-0.05461(6)	0.74984(14)	0.00797(12)	2.38
Ca1	1	0.165020(18)	-0.33992(3)	0.74980(6)	0.00724(11)	2.14
Ca2	1	0.334553(18)	-0.66169(3)	0.24965(6)	0.00702(11)	2.11
Ca3	1	0.030046(18)	-0.87742(3)	0.75001(6)	0.00871(12)	1.82
Ca4	1	0.468770(19)	-0.12255(3)	0.24992(6)	0.00900(12)	1.82
P1	1	0.09954(2)	-0.60150(4)	0.74932(8)	0.00571(14)	5.04
P2	1	0.18753(2)	-0.76951(4)	0.46151(8)	0.00796(15)	5.00
P3	1	-0.07291(2)	-0.76881(4)	0.03647(8)	0.00781(15)	4.96
P4	1	0.31160(2)	-0.23125(4)	0.53685(8)	0.00777(15)	4.98
P5	1	0.60001(2)	-0.60058(4)	0.75046(8)	0.00571(14)	5.05
P6	1	0.57207(2)	-0.23085(4)	0.96156(9)	0.00799(15)	4.97
Al1	0.638(3)	0.084874(18)	-0.65969(3)	0.24922(7)	0.00536(14)	3.07
Fe1	0.362(3)	0.084874(18)	-0.65969(3)	0.24922(7)	0.00536(14)	
Al2	0.877(4)	0.5	-0.5	1	0.0054(2)	3.21
Fe2	0.123(4)	0.5	-0.5	1	0.0054(2)	
Al3	0.873(4)	0	-0.5	0	0.0049(2)	3.20
Fe3	0.127(4)	0	-0.5	0	0.0049(2)	
Al4	0.888(3)	0.24976(3)	-0.50032(4)	0.49943(10)	0.00453(14)	3.19
Fe4	0.112(3)	0.24976(3)	-0.50032(4)	0.49943(10)	0.00453(14)	
Al5	0.637(3)	0.414472(18)	-0.34119(3)	0.74937(7)	0.00512(14)	3.07
Fe5	0.363(3)	0.414472(18)	-0.34119(3)	0.74937(7)	0.00512(14)	
O1	1	0.06526(6)	-0.67687(10)	0.5568(2)	0.0097(4)	1.88
O2	1	0.09631(7)	-0.67656(10)	0.9419(2)	0.0104(4)	1.88
O3	1	0.17750(6)	-0.52316(10)	0.7052(2)	0.0089(4)	2.01

O4	1	0.06060(6)	-0.52346(10)	0.7938(2)	0.0085(4)	2.01
O5	1	0.15321(8)	-0.82621(12)	0.6666(3)	0.0200(5)	1.63
O6	1	0.23182(7)	-0.82658(10)	0.3550(2)	0.0121(4)	1.75
O7	1	0.24255(6)	-0.65035(10)	0.5087(2)	0.0089(4)	1.95
O8	1	0.12772(6)	-0.76739(10)	0.3024(2)	0.0122(4)	1.76
O9	1	-0.14571(7)	-0.82549(10)	0.1453(2)	0.0118(4)	1.80
O10	1	-0.06764(7)	-0.82523(11)	-0.1707(3)	0.0182(5)	1.61
O11	1	-0.01170(7)	-0.76743(10)	0.1953(2)	0.0123(4)	1.74
O12	1	-0.06811(6)	-0.64938(10)	-0.0089(2)	0.0093(4)	1.96
O13	1	0.26682(7)	-0.17526(10)	0.6441(2)	0.0114(4)	1.81
O14	1	0.34527(8)	-0.17442(12)	0.3298(3)	0.0198(5)	1.62
O15	1	0.25679(6)	-0.35064(10)	0.4904(2)	0.0096(4)	1.96
O16	1	0.37185(7)	-0.23305(10)	0.6957(2)	0.0121(4)	1.74
O17	1	0.67767(6)	-0.52231(10)	0.7066(2)	0.0095(4)	2.01
O18	1	0.56105(7)	-0.52258(10)	0.7951(2)	0.0099(4)	2.01
O19	1	0.56563(7)	-0.67586(10)	0.5586(2)	0.0112(4)	1.88
O20	1	0.59669(7)	-0.67590(10)	0.9432(2)	0.0099(4)	1.88
O21	1	0.64489(7)	-0.17392(10)	0.8555(2)	0.0120(4)	1.75
O22	1	0.56598(7)	-0.17416(12)	1.1680(3)	0.0200(5)	1.62
O23	1	0.51111(7)	-0.23324(10)	0.8019(2)	0.0121(4)	1.75
O24	1	0.56752(6)	-0.34989(10)	1.0092(2)	0.0095(4)	1.95
Oh1	1	0.04865(6)	-0.54303(10)	0.2239(2)	0.0088(4)	1.08
Oh2	1	0.45082(6)	-0.45795(9)	0.7770(2)	0.0074(4)	1.08
Oh3	1	0.17966(6)	-0.54293(9)	0.2757(2)	0.0082(4)	1.07
Oh4	1	0.31985(6)	-0.45789(9)	0.7233(2)	0.0080(4)	1.08
Ow1a	0.563(9)	0.20683(17)	0.0663(3)	0.9061(7)	0.0159(6)	0.27
Ow1b	0.412(8)	0.1975(2)	0.0381(4)	0.9666(10)	0.0159(6)	0.32
Ow2a	0.577(9)	0.30248(16)	-0.0399(3)	0.0408(6)	0.0139(6)	0.32
Ow2b	0.393(8)	0.2922(2)	-0.0680(4)	0.1013(10)	0.0139(6)	0.24
Ow3b	0.435(8)	0.32287(19)	0.0388(4)	0.5349(9)	0.0149(6)	0.32
Ow3a	0.569(8)	0.32590(14)	0.0669(3)	0.5950(6)	0.0149(6)	0.27
Ow4a	0.614(9)	0.17772(14)	-0.0401(3)	0.4584(6)	0.0181(6)	0.32
Ow4b	0.417(8)	0.1743(2)	-0.0680(4)	0.3971(9)	0.0181(6)	0.25
Ow5	1	0.54365(7)	-0.05102(10)	0.5802(2)	0.0121(4)	0.42
Ow6	1	0.43122(7)	-0.05171(10)	-0.0808(2)	0.0123(4)	0.43
Ow7	1	0.08081(7)	-0.34501(12)	0.4800(2)	0.0153(5)	0.34
Ow8	1	0.24675(7)	-0.34417(12)	1.0213(2)	0.0171(5)	0.34
Ow9	1	0.25340(7)	-0.65607(12)	-0.0226(2)	0.0171(5)	0.34
Ow10	1	0.41841(8)	-0.65624(13)	0.5227(3)	0.0192(5)	0.34
Ow11	1	-0.04300(7)	-0.94869(10)	0.4190(2)	0.0121(4)	0.43
Ow12	1	0.06904(7)	-0.94975(10)	1.0804(2)	0.0124(4)	0.42
Hoh1	1	0.0692	-0.483	0.303	0.055(2)	
Hoh2	1	0.4284	-0.5207	0.6948	0.055(2)	
Hoh3	1	0.1866	-0.4827	0.2022	0.055(2)	
Hoh4	1	0.3124	-0.5093	0.8041	0.055(2)	
H1a1	0.6	0.1706	-0.001	0.9328	0.055(2)	
H1a2	0.6	0.2188	0.0818	0.7687	0.055(2)	
H2a1	0.6	0.3146	-0.0913	0.1739	0.055(2)	

H2a2	0.6	0.3275	0.0331	0.0968	0.055(2)	
H3a1	0.6	0.3592	0.1007	0.6911	0.055(2)	
H3a2	0.6	0.3321	0.1166	0.5248	0.055(2)	
H4a1	0.6	0.1683	-0.1133	0.4571	0.055(2)	
H5a	1	0.5326	-0.1103	0.6517	0.055(2)	
H5b	1	0.5878	-0.0382	0.5521	0.055(2)	
H6a	1	0.4141	-0.1157	-0.1727	0.055(2)	
H6b	1	0.3925	-0.0367	-0.0511	0.055(2)	
H7a	1	0.0782	-0.2919	0.3856	0.055(2)	
H8a	1	0.2794	-0.2931	1.0973	0.055(2)	
H8b	1	0.2837	-0.3732	0.9477	0.055(2)	
H9a	1	0.229	-0.7139	-0.1001	0.055(2)	
H9b	1	0.2257	-0.6229	0.0635	0.055(2)	
H10a	1	0.4248	-0.7115	0.5761	0.055(2)	
H11a	1	-0.0249	-0.8956	0.3333	0.055(2)	
H11b	1	-0.0885	-0.9697	0.4435	0.055(2)	
H12a	1	0.0845	-0.891	1.1638	0.055(2)	
H12b	1	0.1067	-0.9712	1.0487	0.055(2)	

Table 8. Polyhedral distances for kinglymountite triclinic model for fanfaniite

Mn1-O9	2.0120(13)	Mn2-O6	2.0114(18)
Mn1-O13	2.0151(17)	Mn2-O21	2.0109(14)
Mn1-Ow1a	2.128(5)	Mn2-Ow1b	2.157(6)
Mn1-Ow2a	2.188(4)	Mn2-Ow2b	2.128(7)
Mn1-Ow3a	2.136(3)	Mn2-Ow3b	2.163(5)
Mn1-Ow4a	2.196(4)	Mn2-Ow4b	2.119(4)
av	2.112	av	2.098
Ca1-O3	2.5140(15)	Ca2-O6	2.4368(12)
Ca1-O4	2.5205(12)	Ca2-O7	2.5335(15)
Ca1-O9	2.4304(16)	Ca2-O17	2.5225(15)
Ca1-O12	2.5248(15)	Ca2-O18	2.5257(12)
Ca1-O13	2.4258(12)	Ca2-O21	2.4326(16)
Ca1-O15	2.5290(15)	Ca2-O24	2.5311(15)
Ca1-Ow7	2.3423(16)	Ca2-Ow9	2.3460(16)
Ca1-Ow8	2.3439(16)	Ca2-Ow10	2.3492(17)
av	2.454	av	2.460
Ca3-O1	2.7125(14)	Ca4-O14	2.3708(16)
Ca3-O2	2.7122(13)	Ca4-O19	2.7199(14)
Ca3-O5	2.3718(16)	Ca4-O20	2.7241(13)
Ca3-O10	2.3714(17)	Ca4-O22	2.3667(18)
Ca3-Ow11	2.4750(15)	Ca4-Ow5	2.4882(15)
Ca3-Ow11	2.6160(16)	Ca4-Ow5	2.6050(16)
Ca3-Ow12	2.4929(16)	Ca4-Ow6	2.4738(16)
Ca3-Ow12	2.6010(13)	Ca4-Ow6	2.6165(13)
av	2.544		2.546

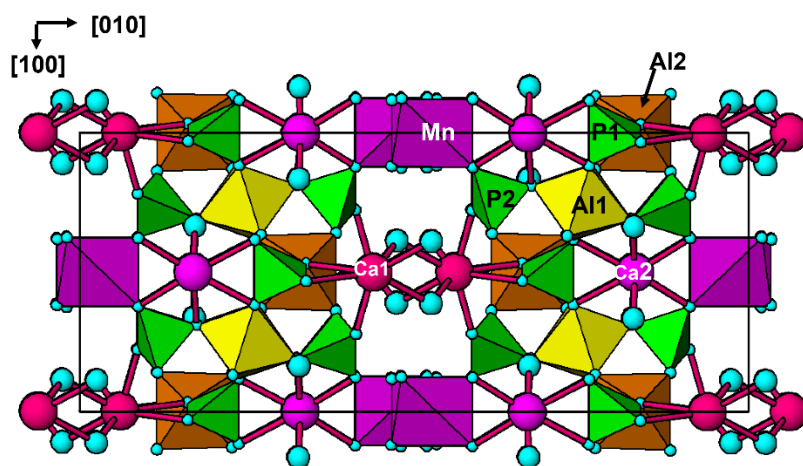
Al1-O1	1.9626(15)	Al2-O18 x2	1.8974(15)
Al1-O2	1.9649(15)	Al2-O24 x2	1.9050(11)
Al1-O8	1.9308(16)	Al2-Oh2 x2	1.8849(14)
Al1-O11	1.9296(12)	av	1.896
Al1-Oh1	1.9269(16)		
Al1-Oh3	1.9324(11)	Al3-O4 x2	1.8980(15)
av	1.941	Al3-O12 x2	1.9010(11)
		Al3-Oh1 x2	1.8883(14)
Al4-O3	1.8968(15)	av	1.896
Al4-O7	1.9042(15)		
Al4-O15	1.9021(15)	Al5-O16	1.9332(16)
Al4-O17	1.8999(15)	Al5-O19	1.9677(15)
Al4-Oh3	1.8861(14)	Al5-O20	1.9648(15)
Al4-Oh4	1.8881(14)	Al5-O23	1.9302(12)
av	1.896	Al5-Oh2	1.9292(15)
		Al5-Oh4	1.9265(11)
P1-O1	1.5323(14)	av	1.942
P1-O2	1.5314(15)		
P1-O3	1.5333(12)	P2-O5	1.5160(16)
P1-O4	1.5328(16)	P2-O6	1.5288(17)
av	1.532	P2-O7	1.5471(12)
		P2-O8	1.5494(15)
P3-O9	1.5313(14)	av	1.535
P3-O10	1.5211(17)		
P3-O11	1.5545(16)	P4-O13	1.5279(17)
P3-O12	1.5466(14)	P4-O14	1.5216(16)
av	1.538	P4-O15	1.5458(12)
		P4-O16	1.5537(15)
P5-O17	1.5321(12)	av	1.537
P5-O18	1.5314(17)		
P5-O19	1.5301(14)	P6-O21	1.5292(14)
P5-O20	1.5321(15)	P6-O22	1.5207(18)
av	1.531	P6-O23	1.5513(15)
		P6-O24	1.5466(15)
		av	1.537

Table 9. H bonding in fanfaniite from Hühnerkobel (Å, °). Kingmountite triclinic model.

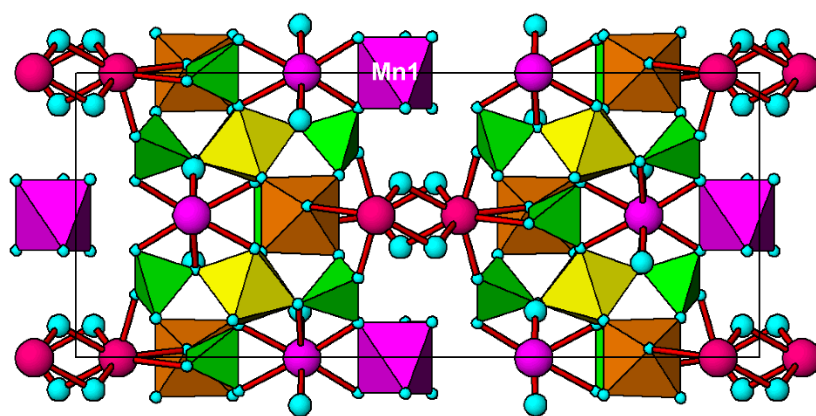
D-H...A	D-H	H...A	D-A	D-H...A
Oh1-Hoh1-Ow7	0.8809(12)	2.0535(16)	2.896(2)	159.75(10)
Oh2-Hoh2-Ow10	0.9214(11)	2.0179(17)	2.892(2)	157.79(10)
Oh3-Hoh3-Ow8	0.8709(13)	2.0853(14)	2.8977(19)	154.93(10)
Oh4-Hoh4-Ow9^	0.8040(13)	2.1249(14)	2.8894(19)	158.87(9)
Ow1a-H1a1-O9	0.916(3)	2.1786(13)	2.915(4)	136.9(3)
Ow1a-H1a2-Ow4a	0.897(4)	2.436(4)	3.075(6)	128.4(2)
Ow2a-H2a1-O14	1.142(4)	1.7229(17)	2.846(4)	166.5(2)
Ow2a-H2a2-O21	0.954(3)	1.7276(13)	2.654(4)	162.6(2)
Ow3a-H3a1-O22	0.872(3)	1.6714(14)	2.531(3)	168.2(2)
Ow3a-H3a2-O21	0.751(4)	2.4843(15)	3.097(4)	140.0(3)
Ow5-H5a-O23	0.8447(14)	1.7588(13)	2.6034(19)	179.02(10)
Ow5-H5b-Ow3a	0.8593(14)	2.136(4)	2.942(4)	156.03(12)
Ow6-H6a-O16	0.9606(13)	1.6513(12)	2.6024(18)	169.86(9)
Ow6-H6b-Ow2a	0.8945(15)	1.895(4)	2.772(4)	166.14(14)
Ow7-H7a-O10	0.9216(16)	2.0918(16)	3.013(2)	177.64(9)
Ow8-H8a-O14	0.8676(13)	2.1515(14)	3.009(2)	169.49(11)
Ow8-H8b-O20	1.0682(17)	2.3204(13)	3.054(2)	124.43(8)
Ow8-H8b-Oh4	1.0682(17)	2.0988(14)	3.110(2)	157.05(8)
Ow9-H9a-O5	0.8687(14)	2.1907(14)	3.032(2)	163.04(12)
Ow9-H9b-Oh3	0.9942(17)	2.1277(14)	3.115(2)	172.01(8)
Ow10-H10a-O22	0.8444(18)	2.2279(16)	3.024(2)	157.29(12)
Ow11-H11a-O11	0.8407(13)	1.8029(14)	2.6022(19)	158.19(12)
Ow11-H11b-Ow4a	0.8644(13)	1.956(3)	2.778(4)	158.25(13)
Ow12-H12a-O8	0.8781(13)	1.7362(13)	2.6062(18)	170.43(10)
Ow12-H12b-Ow1a	0.9268(16)	2.097(4)	2.929(4)	148.69(14)



Fig. 1. Clusters of pale yellow fanfaniite crystals associated with small colourless platelets of paulkerrite-group mineral. FOV 0.6 mm. Image by Christian Rewitzer.



(a)



(b)

Fig. 2. [001] projections of the (a) $C2/c$ average structure and (b) $C2$ structure model for fanfaniite from Hühnerkobel. Only the Mn1 site, with the higher site occupancy is shown.

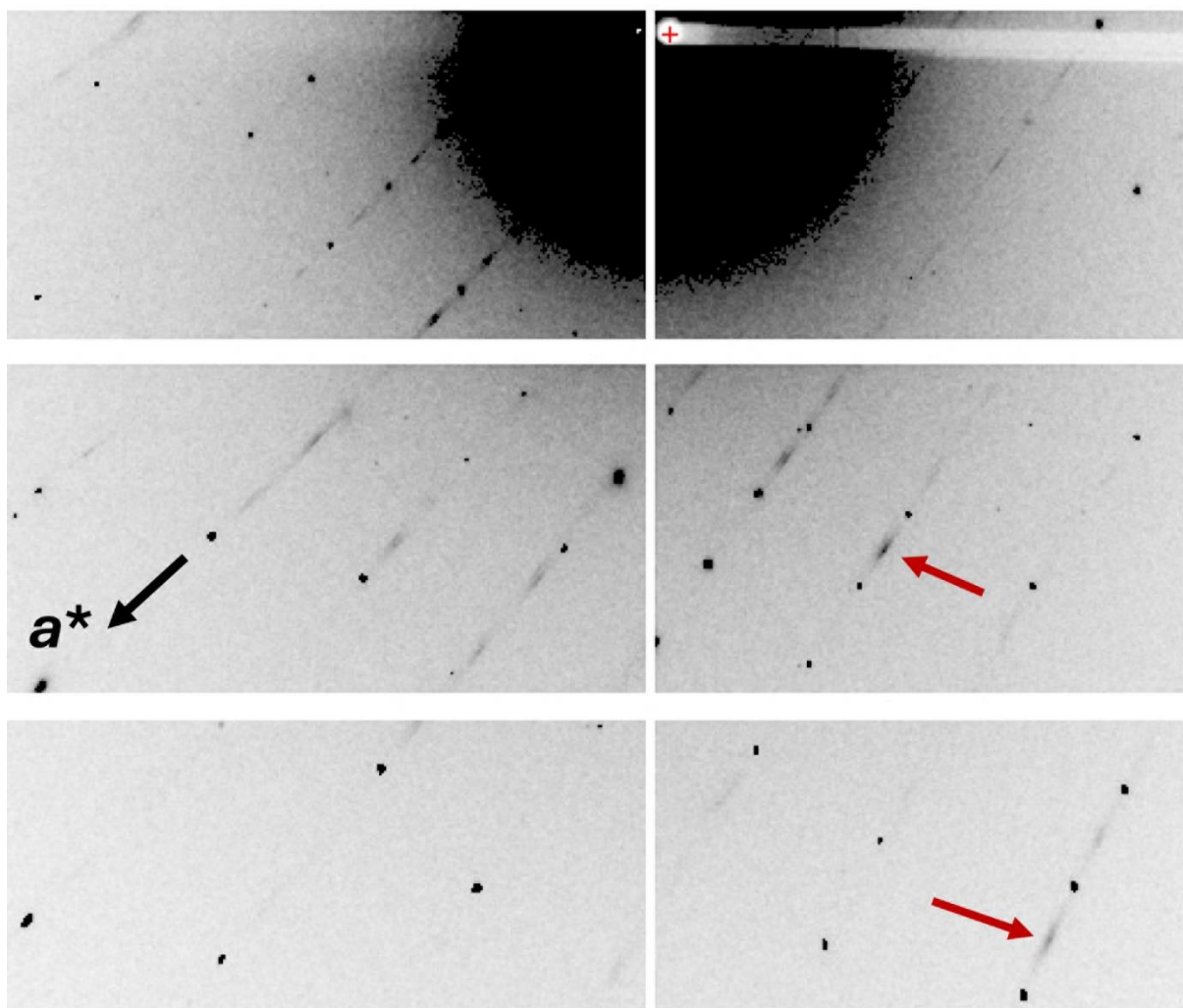


Fig. 3. Screenshot of diffraction by fanfaniite. Red arrows show diffuse reflections corresponding to a doubling of a .

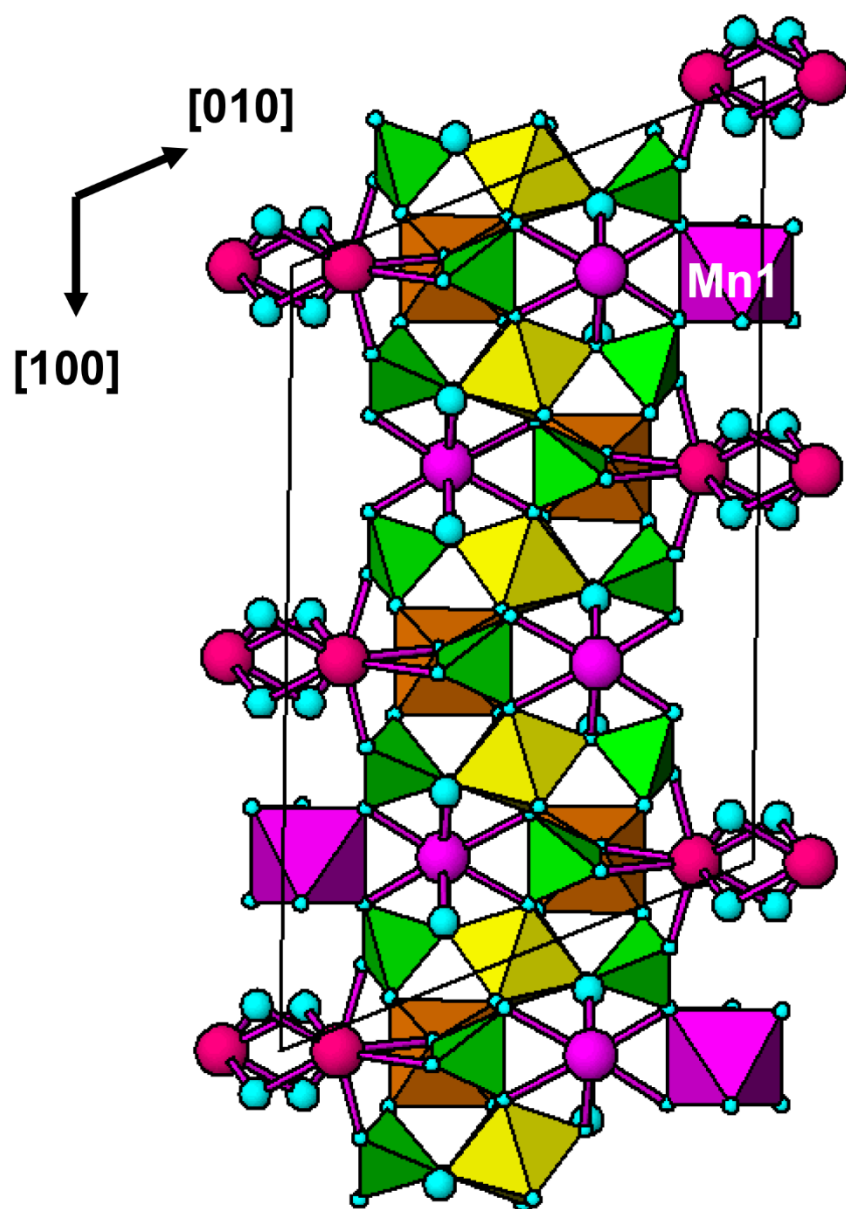


Fig. 4. [001] projection of triclinic *P*-1 model with doubled *a* parameter, for fanfaniite. Only the Mn1 site, with the higher site occupancy is shown.